A New Upper Pliocene Palaeomagnetic Pole from Western Syria and a preliminary Polar Wander Curve of the Arabian Plate

Jamal Abou-Deeb

Department of Geology-Faculty of Sciences-University of Damascus-Syria

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ABSTRACT

44 sites were sampled from the Upper Pliocene (βN_2^b) basaltic flows, from the western flank and the distant eastern flank of the Levant Fault in the Buqeia area NW of Tel Kalakh and west of Homs. Thermal demagnetization led to the identification of consistent directions of remanence in most sites. Normal, Reversed and Intermediate polarities are identified, with Reversed polarities dominating. The Intermediate polarity sites were excluded and the mean directions of the Reversed sites were reversed. The mean site direction of the 16 accepted sites from the western flank of the Levant Fault is 3.6°, 46.5° with α_{95} =7.0°, while the mean site direction of the 10 accepted sites from the eastern flank is 17.5°, 46.8° with α_{95} =2.6°. The overall mean direction of the 26 accepted sites in the western and eastern flanks is 9.0°, 46.9° with α_{95} =4.6°.

While the mean VGP of the 16 accepted sites from the western flank of the Levant Fault is 83.2° N, 188.3° E with $\alpha_{95}=7.0^{\circ}$, and the mean VGP of the 10 accepted sites from the eastern flank is 73.8° N, 145.6° E with $\alpha_{95}=2.6^{\circ}$. The overall mean VGP of the 26 accepted sites in the western and eastern flanks is 80.2° N, 162.4° E with $\alpha_{95}=4.8^{\circ}$.

The entire available palaeomagnetic pole data sources were consulted which resulted in collecting 51 poles from the Arabian plate, ranging in age from Precambrian to Quaternary. These poles were used to construct a preliminary polar wander curve, which will improve with the increase of the future results from Arabia.

Key words: Palaeomagnetism, Levant Fault, Upper Pliocene, Misiaf, Ghab, Syria, Lebanon



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Introduction and Regional Geology

The Levant transform fault system (Ponikarov, 1966; Best et al., 1993) - the northern-most extension of the Dead Sea transform fault system (Dubertret, 1970, Girdler, 1990) enters Syria from northern Lebanon, where it is called the Yemmunah Fault, and, in Syria, forms the Missiaf-Ghab fault system. That zone widens in the north to form the rhombshaped Missiaf-Ghab graben (Sigachev et al., 1995). The Yemmunah and Missiaf-Ghab fault systems are together widely regarded as the major seismic zone and the source of the volcanic activity in the area with a predicted left lateral displacement of about 20-25 km (Quennell, 1984; Trifonov et al., 1991). Although Butler et al. (1997) concluded that neither of these zones had been active during the past 5 Ma but that, during this time, seismic activity had mostly been associated with the Roum Fault that extends from south Lebanon through Beirut. The subaerial flood basalts of the Early Miocene, mostly ranging in age from 20 to 16 Ma ago, extended over most of southern Syria, Jordan and Saudi Arabia, but were absent along the Syrian coastal zone (Mouty et al., 1992). After a period of volcanic activity quiescence lasted till about 8 My ago an intensive volcanism commenced over most of Syria, but particularly along the rift margins in the south and the Levant transform fault margins in the north (Fig. 1). This phase of volcanism remained active into prehistoric time, for example, in the Karasu Valley, Southern Turkey, in the northernmost part of the rift (Capan et al., 1987) where ages range from 2 My to 0.4 My.

The studied area is mainly covered with the Upper Pliocene basaltic flows (βN^b_2) and the Quaternary sequences which are mostly consolidated sediments and are dominated by those of the Al Buqeia depression. The area is divided by the main N/S extending Levant Fault and many other faults, which are mainly trending N-S, NW-SE and NE-SW, but only the main faults are shown in Fig. 1.

Few palaeomagnetic studies were done (Van Dongen et al., (1967), Roperch and Bonhommet, (1986) and Abou-Deeb et al., (1999)). Three areas; 1- Sweida in the southern part of Syria, 2- Kisweh to the south of Damascus and, 3- the Homs-Shin basaltic flows west of Homs, were studied by Abou-Deeb et al., (1999) in order to widen the palaeomagnetic knowledge of the area. The results induced the author to undertake a new study of the Kisweh and the Homs-Shin basaltic flows. The later is the

subject of this study that comprises the results of 44 sites, each comprising 6-7 samples, in Upper Pliocene (βN^b_2) basaltic flows, from the western flank and from the distant eastern flank of the Levant Fault. Generally, three sites were taken at separate locations within similar or identical exposures.



Fig. 1. Outline Geological Map of the studied area.

The sampling localities represent mostly 3 sites per volcanic unit. The main N/S Levant Fault is labelled, but, for clarity, the orientations of only a few relevant faults are marked. The fault trends are mainly N-S, NW-SE and NE-SW directions but only the main faults systems can be shown.

Numbers are as follows: 1=T31-T33, 2=T22-T24, 3=T4-T8, 4=T37-T39, 5=T19-T21, 6=T25-T28, 7=T40-T42, 8=T46-T48, 9=T43-T45, 10=T64-T66, 11=T67-T69, 12=T49-T51, 13=T52-T54, 14=T55-T57.

Sampling, Measurement and Treatment

Palaeomagnetic drilling techniques (Collinson, 1983; Tarling, 1983) were employed by using petrol driven drill and both sun and magnetic compasses were used for orientation. Then cores, 2.5 cm in diameter and 2.1 cm in height, were sliced in the Palaeomagnetic Laboratory of the Geology Department of Damascus University. The magnetometer and demagnetizing instruments of the Geology Department of Plymouth University, England, were used to achieve all the required measurements (Molyneux, 1971, Abou-Deeb, 1997 and Abou-Deeb et al., 1999). Apart from the strongly magnetized T56 site (intensity 116.5 A/m) (Table 1), the initial intensity of magnetization ranged from 31.4 A/m (site T44) to 0.59 A/m (site T66) with an average of 6.3 ± 6.4 A/m. The site low-field susceptibilities ranged from 3.4 mSI (T66) to 71.1 mSI (T22) with an average of 31.8 ± 17.1 A/m. 5 pilot samples from each of the main localities, T7.2, T20.6, T41.5, T54.2, T67.4, being treated at 100, 150, 200, 250, 300, 350, 400, 450, 500, 520, 540, 560 & 580°C. Sample characteristic directions were defined where the vectors isolated had a mean angular deviation (MAD) of $< 5^{\circ}$ (Kirschvink, 1980). A site mean direction was considered further if it was defined by at least 5 samples with characteristic directions that agreed with an estimated 95% Probability cone of confidence of $<15^{\circ}$ (Fisher, 1953), that is the cone angle which is characterized by 95% probability that the true direction is located inside the cone is less than 15°. Alternating magnetic field demagnetization up to 100 mT was also undertaken on 5 samples from the same representative sites, but no reliable vectors could be isolated in any of these samples by using this technique. According to the study of the behavior of the pilot samples the rest of the samples were thermally demagnetized at 200, 300, 350, 400, & 420 °C.

The magnetic susceptibility (Fig. 2) and intensity of magnetization (Fig. 3) of the pilot sample were studied in order to determine if any thermo-chemical changes had occurred to the magnetic minerals during heating. This will help in the extrapolation of the magnetic mineral that carries the magnetization of the rocks.

Fig. 2 shows variable changes in the magnetic susceptibility of all samples which indicate the presence of small amounts of other magnetic

minerals and started to change, after heating to $50 \,^{\circ}\text{C}$ with a major change started at 250 °C, to other minerals of higher susceptibility (like magnetite). While the decrease in the magnetic susceptibility after heating between 500 °C and 580 °C probably due to oxidation of magnetite to haematite.



Fig. 2. Curves of normalized magnetic susceptibility of the pilot samples. Sample numbers and symbols are shown on the left side of figure.

In spite of the continuous decrease of the magnetic intensity of sample T54.2, the normalized magnetic intensity of all samples (Fig. 3) suggests that the main carrier of magnetization is magnetite.



Fig. 3. Curves of normalized intensity of magnetization of the pilot samples. Sample numbers and symbols are shown on the left side of figure.

Results and Interpretations

In the majority of sites, the samples had well defined characteristic directions during thermal demagnetization, with their mean angular deviations (*MAD*) mostly being less than 3°. The main exceptions were in 2 sites (T4, T8) in which most samples *MAD*'s exceeded 5°; these sites were excluded from further directional analyses. In the remaining sites, the blocking temperature (Fig. 3) spectra suggested that the characteristic magnetic directions were carried by magnetite. Above about 250°C, the low field susceptibility tended to increase (Fig. 2), but with no apparent effect on the linearity of the vector until approaching the Curie point of magnetite. A well defined site mean direction was isolated in most remaining sites, with the site mean precision estimate, α_{95} (Fisher, 1953), being <15° (Table 1). However, 1 site was rejected (T40) in which $\alpha_{95}>15^\circ$.

Normal (N), Reversed (R) and Intermediate (I) polarities were present (Table 1). According to the definitions (N = site virtual geomagnetic pole [VGP] latitude >50°N; R = site VGP >50°S; I = site VGP between 50°N and 50°S), the majority of the sites, 23, were of Reversed polarity, 4 were of Normal polarity and 14 were of Intermediate polarity. The mean site initial intensity of magnetization of the 23 Reversed polarity sites $(3.9\pm3.1 \text{ A/m})$ was almost similar to that of the 4 Normally magnetised sites $(4.1 \pm 1.1 \text{ A/m})$. The 14 sites of Intermediate polarity were more strongly magnetised and much more scattered (11.5±9.0 A/m) even after the removal of the most strongly magnetised site (T56). While the low-field susceptibility site mean values were almost similar, irrespective of polarity: Reversed (31.0±19.9 mSI), Normal (40.2±5.9 mSI), Intermediate (31.8 ± 13.6 mSI).

For calculating the VGP, all 14 sites of Intermediate polarity were ignored and site T38 was excluded because of its very shallow inclination (7.7°), which could have been caused by local movement and the polarity of the remaining 22 Reversed sites were reversed by 180°. Then the remaining 26 mean site vectors were northerly directed and have positive inclination. The mean site direction of the 16 accepted sites (T19-T24, T27, T31-T33, T37, T39, T42, T46-T48) from the western flank of the Levant Fault is 3.6°, 46.5° with α_{95} =7.0° (Table 2), while the mean site direction of the 10 accepted sites (T49-T51, T57, T64-T69) from the eastern flank is 17.5°, 46.8° with α_{95} =2.6° (Table 2). The overall mean

direction of the 26 accepted sites from the western and eastern flanks is 9.0°, 46.9° with α_{95} =4.6°. Such a direction is consistent with that (5.6°, 49.9°), predicted for the Arabian plate some 10 Ma ago, after allowing for the opening of the Red Sea and using the African polar wander curve (Besse & Courtillot, 1991). This implies that the studied locality belongs to unrotated part of the Arabian Plate.



Fig. 4. Site mean directions after demagnetization plotted on an equal area projection. Positive (downward) inclinations are shown with solid circles and negative (upward) inclinations are shown with hollow circles. The solid triangle is the mean of the Positive inclinations and the (x) sign is the mean of the negative inclinations. The centre of the hollow inverted triangle is the overall mean site direction of the accepted sites after reversing the negative inclinations.



Fig. 5. Virtual Geomagnetic Poles derived from site mean directions plotted in Fig.4. The overall mean of the VGP's is represented by the centre of the (X) sign.

While the mean VGP of the 16 accepted sites from the western flank of the Levant Fault is 83.2° N, 188.3° E with α_{95} =7.0° (Table 2), and the mean VGP of the 10 accepted sites from the eastern flank is 73.8° N, 145.6° with α_{95} =2.6° (Table 2). The overall mean VGP of the 26 accepted sites from the western and eastern flanks is 80.2° N, 162.4° with α_{95} =4.8°

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Table 1. Site Mean Intensity, Susceptibility, Direction of Remanence and pole position. The number of samples per site, with a vector defined with a mean angular deviation (*MAD*) less than 5° (Kirshvink, 1980) is N, compared with the total number of samples collected, n. The mean initial intensity of magnetization (Int.) is in units of A/m and the mean susceptibility (Susc.) are in mSI units. The site mean directions (Declination, Decl. and Inclination, Incl.) are listed with the estimates of precision, *k*, and 95% Probability, a_{95} (Fisher, 1953). The VGP latitude (Lat.) and longitude (long.) are calculated for each site. The polarity (Pol.) of each site is Normal (N), when the virtual geomagnetic pole latitude is greater than 50°N, Reversed (R), when greater than 50°S, and Intermediate (I) when between these two limits. "Scattered" sites (S) are when there are fewer than 5 samples with *MAD* < 5° or when the site precision, a_{95} , exceeds 15°.

Site	N/n	Int. A/m	Susc. mSI	Decl.	Incl.	k	α_{95}	Lat.	Long. I	Pol.
Easter	n Flai	nk								
Ain-K	ut									

	Г64	6/6	1.69±0.37	13.86±4.98	199.4	-43.1	1915	1.5	-70.6	331.3	R
	Г65	6/6	1.03±0.12	6.17±0.43	199.0	-45.9	1760	1.6	-72.1	325.9	R
	Г66	6/6	0.59±0.06	3.44±0.34	194.2	-50.2	727	2.5	-77.5	320.4	R
	Г67	5/6	1.58±0.26	27.17±3.60	199.4	-51.6	574	3.2	-73.7	309.9	R
	Г68	6/6	1.60±0.23	24.87 ± 4.98	202.4	-40.8	108	6.5	-67.4	330.8	R
	Г69	6/6	2.46±0.68	40.33±2.26	199.7	-44.9	440	3.2	-71.2	327.2	R
Suf	fer										
	Г49	6/6	3.86±0.41	33.52±2.04	18.1	46.9	618	2.7	73.2	144.9	Ν
	Г50	6/6	3.83±0.42	40.63±2.90	12.2	45.8	379	3.4	77.1	159.3	Ν
	Г51	6/6	3.13±0.30	47.90±1.77	12.2	46.5	643	2.6	77.5	157.3	Ν
	Г52	6/6	11.87 ± 3.80	48.43±1.82	314.7	-1.7	659	2.6	34.7	276.3	Ι
	Г53	6/6	19.55±3.20	50.98±6.23	308.8	0.2	452	3.2	31.0	281.9	Ι
	Г54	5/6	20.15±0.98	39.81±1.71	254.7	-1.2	1126	2.3	-12.9	314.9	Ι
	Г55	6/6	4.95±1.70	40.66±3.95	256.1	-8.0	112	64	-13.7	311.1	Ι
	Г56	6/6	116.50±5.09	37.36±2.49	257.5	-17.5	325	3.7	-15.3	306.1	Ι
	Г57	5/6	5.75±1.95	38.91±6.65	17.8	52.4	122	6.9	75.1	128.3	Ν
We	ester	n fla	nk								
Ma	rma	rita									

T40	5/6	7.43±1.08	36.58±2.27	Scattered		21	17.0	scattered		S
T41	5/6	5.30±0.19	8.61±1.76	112.7	-62.2	1671	1.9	-38.5	157.5	Ι
T42	6/6	2.20±0.80	6.83±1.12	170.5	-66.9	134	5.8	-73.7	193.8	R
T43	6/6	21.47±3.29	30.87±1.45	73.9	4.3	161	5.3	14.4	133.9	Ι
T44	6/6	31.38±4.87	28.09±0.97	57.4	5.4	423	3.3	28.0	144.0	Ι
T45	6/6	12.20±1.47	21.89±3.58	33.4	16.3	300	3.9	49.6	159.3	Ι
T46	5/6	1.86±0.45	63.40±2.14	175.8	-54.1	165	6.0	-86.5	125.2	R
T47	5/6	2.27±0.26	49.71±1.52	182.5	-43.1	318	4.3	-80.0	22.9	R
T48	6/6	1.87±0.17	53.61±2.18	185.4	-42.4	483	3.1	-78.7	10.5	R
O-let I	71 II									

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T19	6/6	5.18±0.32	21.20±1.67	195.5	-49.2	944	2.2	-76.2	321.9	R
T20	5/6	2.94±0.10	6.75±0.71	189.9	-49.9	1319	2.1	-80.8	329.9	R
T21	6/6	3.09±0.14	10.57 ± 2.00	187.7	-48.1	1375	1.8	-81.4	344.8	R
T25	4/7	4.70±1.36	37.96±3.48	132.1	-18.3	43	14.2	-39.5	107.9	Ι
T26	6/6	4.39±0.64	44.10±2.97	131.0	-13.2	308	3.8	-37.0	106.1	Ι
T27	6/6	2.86±0.86	32.05±1.34	145.0	-24.1	178	5	-51.4	100.0	R
T37	6/7	8.37±1.63	20.00±6.22	192.8	-32.5	162	4.8	-69.4	359.3	R
T38	5/6	12.41±1.82	22.46±2.40	209.0	-7.7	54	10.5	-49.1	348.7	R
T39	6/6	11.34±1.86	22.69±2.01	195.6	-21.2	240	4.5	-62.3	1.6	R
Tel Ka	lakh									
T4		2.81±0.48	32.42±0.99	Undefined				undefined		S
T5	5/6	5.40±1.89	31.87±6.74	120.3	-59.9	121	7.0	-43.3	152.3	Ι
T6	4/6	3.41±1.40	11.32±5.11	109.6	-56.1	175	7.0	-34.1	150.4	Ι
T7	5/6	4.20±0.67	13.39±2.63	115.6	-54.7	231	5.0	-38.2	146.6	Ι
T8		4.45±0.43	7.86±0.44	undefined				undefined		S
T22	6/6	2.73±0.36	71.05±3.01	175.8	-48.6	467	3.1	-83.7	72.2	R
T23	6/6	3.96±1.45	57.19 ± 11.02	177.5	-51.4	1057	2.1	-86.6	75.8	R
T24	6/6	5.93±1.04	65.42±4.41	173.7	-49.6	1091	2.0	-83.2	89.6	R
T31	6/6	3.76±0.27	30.20±2.18	190.9	-51.2	77	7.7	-80.4	320.4	R
T32	7/7	3.86±0.96	33.46±1.77	197.5	-47.5	347	3.2	-73.9	323.6	R
T33	7/7	5.54±1.16	31.71±0.78	198.2	-49.6	45	9.1	-74.1	316.8	R

 Table 2. The mean directions and VGP positions and parameters of the accepted sites in the western and eastern flanks of the Levant Fault after reversing the negative inclinations. Symbols as in Table 1.

reversing the negative menhations. Symbols as in Table 1.										
No of Site	No. on	Decl.	Incl.	k	α95	Lat.	Long.	α_{95}	Pol.	
	samples									

Mean of the accepted sites in the western flank (T19-T24, T27, T31-T33, T37, T39, T42, T46-T48)

10	90	3.0	40.5	29	/.0	83.2	188.3	7.0	Reverse
N C (1		· · · · · · · · · · · · · · · · · · ·		M 1	(T 10 5	TC1 TC'	7 77 (4 7		

-

 Mean of the accepted sites in the eastern flank
 (T49-T51, T57, T64-T69)

 10
 58
 17.5
 46.8
 352
 2.6
 73.75
 145.6
 2.6
 Mixed

Overall mean of the accepted sites (T19-T24, T27, T31-T33, T37, T39, T42, T46-T51, T57, T64-T69)

26 154 9.0 46.9 40 4.6 80.2 162.4 4.8 Mix

Arabian plate palaeomagnetic poles

All the available palaeomagnetic pole data sources were consulted (The Norwegian Geological Society Internet pole listing, Palaeomagnetic directions and pole positions compiled by M.W. McElhinny (1968, 1971, 1972, 1977), palaeomagnetic database compiled by J. Piper (1988), Palaeomagnetism and Plate tectonic written by M.W. McElhinny (1973) and all the available papers) in order to collect all the data. This resulted in collecting 51 poles, including the present result, (Table 3) from the Arabian plate, and ranging in age from Precambrian to Quaternary. No strict criterion on the validity of the data was applied due to the low number of the geographically widely dispersed studies, with the hope that the future increase of the palaeomagnetic studies will enforce such criterion in order to get a more reliable polar wander curve for the Arabian plate.

Table 4 gives the pole means of the different geological ages. Quaternary (Q) and Pliocene-Pleistocene (T_p - Q_p) poles resulted from one pole each. While Pliocene to Lower Cretaceous poles were chosen according to good grouping of poles i.e. ($\alpha_{95} < 20^\circ$). Also Ordovician to Upper Cambrian (O – ε u), Cambrian (ε), and two Lower Cambrian (ε I) and Precambrian (PC) poles resulted from one pole. Those poles which resulted from only one pole were used separately due to the difference in absolute age. Fig. 6 shows the polar wander curve of the poles shown in Table 4, but the Precambrian pole was not drawn because it looks very strange, which may have resulted from local movement.



Fig. 6. Polar wander curve of the Arabian plate. Pole numbers are as in column two of Table 4. The hollow circles are on the southern hemisphere

Table 3. Summary of all the available Arabian plate poles. Symbols are: Si = No. of sites, Sa = No. of samples, α_{95} = is the cone angle which is characterized by 95% probability that the true direction is located inside the cone. dp, dm=error parameters on the pole position. t = thermal treatment, a = a. f. current treatment. Q=Quaternary, Qp= Pleistocene, Tp = Pliocene, Tm=Miocene, To=Oligocene, Te=Eocene, Ku=Upper Cretaceous, Kl=Lower Cretaceous, J=Jurassic, Ju=Upper Jurassic, Jl=Lower Jurassic, Tr=Triassic, P=Permian, O=Ordovician, E=Cambrian, Eu=Upper Cambrian, El=Lower Cambrian, PC = Precambrian. The top symbol or numbers in column 3 are the geological age of rock while the bottom number is the age of magnetization.

Pole	Rock unit and Coordinates	Age	Si (Sa) &	Pole Po	osition	α	Polarit	Reference
No.		(My)	treatment	Lat.	Long.	α_{95}	У	
						or dn. dm		
1	Holocene Basalts Sweida	0	14 (86) t	85.6N	204 F	11 6	М	Abou-Deeb et al
1	Svria 32.6 N. 36.3 E	×	11(00)1	05.01	20.1 L	11.0		(1999)
2	Pleistocene basalt lava, Jordan	Tp-Op	4 (25) a	73.8N	241.3	4.9	М	Sallomy and Krs
	31.43 N, 35.79 É	0-2			Е			(1980)
3	Volcanics, group1, Palestine	Tp-Qp	5(30) a	77.7 N	200.1	20.4	М	Nur & Helsley
	33.4 N, 35.5 E	0-5			E			(1971)
4	Pliocene basalts, Lebanon	Тр		83.4 N	203.6	7.7		Gregore et al.
		2-4			E			(1974)
5	Upper Pliocene basalts,	Тр	15(78) a	88 N	169 E	7.7	Μ	Gregore et al.
	Lebanon 33.4 N, 35.8 E	2-4						(1974)
6	Hatrurim formation, Palestine	Тр	18 (119) a	82.9 N	135.6	5.2		Ron & Kolodny
	31.5 N, 35.3 E	2-5			E			(1992)
7	Volcanics, group2, Palestine	Тр	3 (46) a	32 N	313.0	13.0	М	Nur & Helsley
0	33.5 N, 35.5 E	2-15	2 (7)	73 0	E		n	(1971)
8	Syrian Basalts, Syria	Tp	2(7)	73.0	251.0	10.2	R	Van Dongen et
0	34.8 N, 35.8 E	2-5	11()	05.11	E 200 F	10,3		al. (1967)
9	Aden Volcanics, Yemen	Ip	11(-) a, r	85 N	298 E		М	l arling et al.
10	13.0 N, 43.0 E) Tu	2((154)+	90 2 M	1(2.4	4.0	м	(1907) December 1-1
10	U-Phocene Basalts, W. Of	1p	20 (154) t	80.2 N	162.4 E	4.8	IVI	Present study
11	Homs, Syna 54.7 N, 50.5 E	J-0 Tn	0(52) +	90.2 M	E 105.9	11.7	м	Ahay Daah at al
11	Homs Syria 34.4 N 36.2 F	1p 5-6	9 (32) t	80.3 N	195.8 F	11./	IVI	(1999)
12	Little Aden Volcanics, Yemen	Tn	5(16) a	78 N	147 E		R	Tarling et al
12	13.0 N. 45.0 E	5-6	5(10) u	,010	11, 2			(1967)
13	Yemen Volcanics, Yemen	Tm	33 (124) a	83.5 N	292.3	6.3		Tarling et al.
_	13.0 N, 45.0 E	5-10	().		E			(1967)
14	Aden Volcanics, Yemen	Tm	12 (54) a	83.0 N	310.0	3.0		Irving & Tarling
	12.8 N, 45.0 E	5-10			Е			(1961)

15	Neocene Basalts, Jordan 32.2 N, 36.5 E	Tm 5-23	1 (6) a	76.2 N	99.1 E	2.1		Sallomy and Krs (1980)
16	Jabel Khariz Volcanics, Yemen 13.0 N, 45.0 E	Tm 10	14(47) a, r	77 N	309 E		М	Tarling et al. (1967)
17	Mount Timna igneous rocks, Palestine 29.8 N, 34.9 E	610-620 10-20	6 (79) a, r	83.6 N	223.2 E	4.4		Marcos et al. (1993)
18	Middle Miocene basalts, Kisweh, Syria 33.2 N, 36.1 E	Tm 16-20	23 (134) t	74.6N	196.0 E	4.4	М	Abou-Deeb et al. (1999)
19	Basalts, Saudi Arabia	Tm 18-20		86.3 N	199.5 E	5.2		Yousif & Beckmann (1981)
20	Miocene Volcanics, Syria 33.25 N, 36.25 E	Tm 19.5	23(108) a, r	75 N	220 E	5,9	М	Roperch and Bonhommet (1986)
21	Hudyadah Volcanics, Yemen 15.0 N, 44.0 E	Tm 23.3	7(38) t	78.5 N	182.8 E	12.5	N	Abou-Deeb (2004)
22	Masirah ophiolite, Oman 20.4 N, 58.8 E	145-152 20-40	-(18) t	78.0 N	240.0 E	5.9	N	Gnos & Perrin (1996)
23	Upper lavas, Madrakah, Saudi Arabia	То 24-30		80.9 N	247.9 E	10.1		Yousif & Beckmann (1981)
24	Sana'a Volcanics, Yemen 15.2 N, 44.1 E	То 29.6	8(46) t	76.6 N	251.3 E	9.4	R	Abou-Deeb et al. (2002)
25	As Sarat Volcanics, Saudi Arabia	To-Tm 29-31	42 (-)	78.8 N	247.8 E	4.3	М	Kellogg & Reynolds (1983)
26	Lower lavas, Madrakah, Saudi Arabia	Те 35-45		68 N	242.8 E	4.2		Yousif & Beckmann (1981)
27	Masirah ophiolite, Oman 20.4 N, 58.8 E	145-152 50-60	-(15) t	56.0 N	282.0 E	8.5	М	Gnos & Perrin (1996)
28	Usfan formation, Saudi Arabia	Kl 65-74		62.9 N	231.2 E	13.2		Yousif & Beckmann (1981)
29	Cretaceous Lavas (1), Palestine 32.5 N, 35.0 E	Ku 65-89	2(15) a	41.6 N	264.2 E	5.4	N	Helsley & Nur (1970)
30	Cretaceous sediments, Palestine 31.0 N, 35.0 E	K 65-146	4() a	77.1 N	250.8 E	19.6		Freund & Tarling (1979)
31	Oman ophiolite unit V3, Oman 24.2 N, 53.6 E	K 74-87	-(27) a	66.4 N	231.7 E	3.7	N	Perrin et al. (1994)
32	Oman ophiolite unit V2, Oman 24.2 N, 53.5 E	K 88-94	-(18) a	11.0 N	314.4 E	4.9	N	Perrin et al. (1994)

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33	Wadi Kadir gabbro, Oman 22 8 N 58 6 E	K 93-98	16 (45) a	60.5 N	285.0 E	7.6	N	Luyendyk & Day (1982)
34	Sheeted dyke, Samail ophiolite, Oman 22.8 N, 58.6 E	K 93-98	28 (202) a	75.6 N	256.5 E	7.0	М	Luyendyk et al. (1982)
35	Oman ophiolite unit V1, Oman 24.2 N, 53.5 E	К 94-105	-(35) a	41.8 N	284.4 E	2.7	N	Perrin et al. (1994)
36	Cretaceous Lavas (2), Palestine 31.0 N, 35.0 E	Kl 97-146	3(22) a	52.8 N	265.1 E	8.6	N	Helsley & Nur (1970)
37	Volcanics and sediments, Lebanon 34.0 N, 36.0 E	Kl 97-146	5(15) a	38 N	282 E	8.9	N	Van Dongen et al. (1967)
38	Essexite Laccolith, Palestine 30.6 N, 34.9 E	J- K 97-157	3() a	48.3 N	265.6 E	6.8, 13.6	N	Freund & Tarling (1979)
39	Volcanics, Mount Harmon, Lebanon	Kl 113-119		4.3 N	301.5 E	15.5		Ron (1987)
40	Neocomian basalts, Lebanon 34.0 N, 35.6 E	Kl 132-146	9(47) a	25 N	285.0 E	9.0	М	Gregore et al. (1974)
41	Masirah ophiolite, Oman 20.4 N, 58.8 E	J 145-152	-(29) t	3 N	271.0 E	9.7	N	Gnos & Perrin (1996)
42	Kimeridgian basalts, Lebanon 34.0 N, 35.8 E	Ju 146-156	8(52) a	2 N	114 E	10.6	R	Gregore et al. (1974)
43	Basalts and tuffs, Lebanon 34.1 N, 35.8 E	Ju 152-155	6(20) a	1 N	120 E	4.3	R	Van Dongen et al. (1967)
44	Ardon formation, Palestine 30.6 N, 34.8 E	Ju 157-208	5() a	74.8 N	248.7 E	15.6	М	Freund & Tarling (1979)
45	Saharonim formation, Palestine 30.6 N, 34.9 E	Tr 208-245	3() a	48.6 N	179.7 E			Freund & Tarling (1979)
46	SHDI sediments, Saudi Arabia 24.0 N, 45.0 E	P - Tr 241-248	- (16) t	53.2 N	250.2 E	15.9		Torq et al. (1997)
47	Red sandstones, Jordan 29.7 N, 35.3 E	O - Eu 470-505	(20) a, t	37 N	323 E	7.5	М	Burek (1969)
48	Diorite Stock, Saudi Arabia 21.7 N, 43.7 E	€ 505-545	1 (12) a, t	25.8 N	332.2 E	11.8		Kellogg & Beckmann (1983)
49	Dykes, Jordan 29.5 N, 35.1 E	El 525-545	2(12) t	26.0 S	341.0 E	8.6		Sallomy and Krs (1980)
50	Mirbat sandstone, Oman 17.1 N, 54.8 E	El 530-560	2 (10) t	23.3 S	321.8 E	7.2	N	Kempf et al. (2000)
51	Arfan Formation, Saudi Arabia 21.3 N, 43.7 E	PC 590-620	12 (63) a, t	77.4 N	297.9 E	10.8		Kellogg & Beckmann (1983)

Table 4. Gives the mean of the Arabian poles given in Table 3. Column one givesthe age as in Table 3, column two gives the pole number on Fig. 6 andcolumn three gives the numbers of the used poles to calculate the mean asgiven in Table 3.

Age	Pole number on Fig. 6	Used Poles	Lat. (N)	Long. (E)	$lpha_{95}$
Q	1	1	85.6	20.4	11.6
Tp-Qp	2	2	73.8	241.3	4.9
Tp-Qp	3	3	77.7	200.1	20.4
Тр	4	4-6, 8-12	84.2	193.5	5.7
Tm	5	13-21	85.2	220.4	6.7
To-Tm	6	22-26	76.5	245.5	4.8
Ku	7	27-31, 33-34	64.4	260.2	11.8
Kl	8	36-38,40	41.4	275.8	16.8
O - Eu	9	47	37.0	323.0	7.5
e	10	48	25.8	332.2	11.8
El	11	49	26.0 S	341.0	8.6
El	12	50	23.3 S	321.8	7.2
PC	13	51	77.4	297.9	10.8

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